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1 **Statistical variability of the correlation plasticity index versus**
2 **liquid limit for smectite and kaolinite**

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15 **ABSTRACT**

16 An extensive literature review has been conducted to observe the statistical correlation of the
17 plasticity index, PI, with the liquid limit, LL, of smectite and kaolinite. Fifty-nine data for
18 smectite and fifty-one for kaolinite have been plotted and compared to each other. The results
19 show that PI is related to LL with equations $PI = 0.97 \times LL - 37.6$ for smectite and $PI =$
20 $5.94e^{0.023 \cdot LL}$ for kaolinite. An independent data set was used for the validation of the proposed
21 relationships. Besides, it was possible to identify a confidence interval for PI, relative to a
22 certain interval for LL values, to confirm the robustness of the relations given above. The
23 findings of this research show that the relation between the Atterberg limits is clearly
24 controlled by the clay mineralogy and that there is no unique way to get PI from LL if the
25 clay mineralogy is not considered.

26 **Keywords:** Atterberg limits; smectite; kaolinite; probabilistic analysis; confidence interval.

27 NOTATION LIST

28	PI	Plasticity index (%)
29	LL	Liquid limit (%)
30	PL	Plastic limit (%)
31	R^2	Regression coefficient
32	k	Number of the available experimental points
33	\bar{m}	Mean values of m (linear regression equation)
34	\bar{n}	Mean value of n (linear regression equation)
35	$t_{\eta\%,n-2}$	Parameter t of Student, relative to a confidence level of $\eta\%$ and a degree of
36		freedom of $k-2$
37	σ	Standard deviation
38		

39 INTRODUCTION

40 Liquid (LL) and plastic limits (PL) are the basic geotechnical index parameters for the
41 qualitative assessment of the physical properties of fine-grained soils. Atterberg (1911), who
42 described first these parameters, stated that “*the liquid limit represents the state at which two*
43 *small pieces of clay placed in a bowl no longer flow together when a bowl is struck violently*
44 *and repeatedly on the hand*” (Haigh, 2012), and it physically describes the water content at
45 the transition from the liquid to the pulpy state of a soil. Soils at LL have small shear strength,
46 which is in the range between 0.5–5.6kPa (e.g. Wasti & Bezirci, 1986; Sridharan & Prakash,
47 1998). PL represents the water content at the transition from the rigid to the semi-solid state
48 of a soil (DIN, 1997). Currently, there are two methods for obtaining LL: the Casagrande
49 (1932) standardized cup method, which is a procedure currently defined in DIN 18122 part 1
50 (1997), AASHTO T89-07 (2007) and ASTM D4318-10 (2010), and the fall-cone-based
51 method, standardized by the ISO/TS 17892-12 (2004) (Spagnoli, 2012). Regarding PL, the
52 geotechnical standard procedure is given by the rolling test method.

53 Two different clay types have been investigated which represent the two extreme types of
54 clay minerals: kaolinite (2-layered clay minerals) and smectite (3-layer clay minerals).

55 The latter has a 2:1 silica:alumina structure, with very weak van der Waals’ forces (Sridharan,
56 2014), and repulsive forces between clay platelets which govern swelling (Taylor and Smith,
57 1986), mainly for Na-smectite (Olson and Mesri, 1970).

58 For kaolinite, positive cation exchange capacity was measured under low pH conditions when
59 edges are positively charged indicates that some isomorphous substitution must exist
60 (Mitchell and Soga, 2005) which was also shown by Brady et al. (1996) and Israelachvili
61 (2011).

62 As stated by Sridharan and Venkatappa Rao (1975), as these clays represent the extreme
63 types of clay minerals, any natural clay is likely to behave, from the geotechnical point of

64 view, in between these two. It is important to observe how the Atterberg limits will change
65 depending on the clay mineralogy. Several authors tried already to assess the behavior of
66 clays regarding their basic geotechnical properties (e.g.: White, 1949; Seed et al., 1964;
67 Dusseault and Scafe, 1979; Nagaraj and Jayadeva, 1983; Sivapullaiah and Sridharan, 1985;
68 Sridharan et al., 1988; Panadian & Nagaraj, 1990; Mitchell and Soga, 2005; Polidori, 2007;
69 Dolinar & Škrabl, 2013), however no unique correlation was given. White (1949) found that
70 LL of kaolinite increased with decreasing particle size, whilst Seed et al. (1964) obtained a
71 linear correlation between LL and percentage of clay size for washed sand with kaolinite.
72 Nagaraj & Jayadeva (1983) found a relationship, where the plasticity index, PI, was $0.74 \times$
73 $(LL-8)$, based on statistical approaches, critical state concepts and on the Gouy-Chapman
74 theory of double layer. However, as stated by Sridharan (2014), since kaolinite and smectite
75 behave quite differently from each other, the mechanisms governing the Atterberg limits, and
76 in turn LL, of kaolinite and smectite are different. The present study summarizes the results
77 of forty-four published data, where the Atterberg limits for almost pure clays were given.
78 From the statistical point of view the confidence interval for confidence levels of 95 and 99%
79 for both smectite and kaolinite has been assessed.

80 **METHODOLOGY**

81 Data from literature about LL and PL for smectite and kaolinite (or well-known natural clays
82 with a predominant clay mineralogy) were carefully analyzed. Only Atterberg limits obtained
83 with the Casagrande cup and the rolling method were used. Regarding smectitic clays, where
84 possible, the main cation was indicated. However, according to Bain (1971), it is possible to
85 roughly distinguish between Na-smectite and Ca-smectite considering their PI values. Clay
86 fraction less than $2\mu\text{m}$ were also indicated. Only Foreman and Daniel (1986), indicated the
87 clay fraction corresponding to $5\mu\text{m}$.

88 Forty-four different published data were used to obtain the Atterberg limits for pure clays, i.e.
89 smectitite and kaolinite mixed with water. Tab. 1 and 2 summarize the values used for the
90 interpretation of the Atterberg limits for natural clays mixed with water. LL values are
91 obtained with the Casagrande cup. Fifty-nine data were used for the interpretation for the
92 smectite, and fifty-one for the kaolinite.

No.	Reference	< 2 μ m (%)	LL (%)	PL (%)	PI (%)	Cation
1	Seed et al. (1964)	95.5	521.5	48	473.5	not specified
2	Andrews et al. (1967)		506	55	451	Na
3	Mesri and Olson (1971)	97	675	39	636	not specified
4	Sridharan et al. (1973)	-	305	44	261	not specified
5	Sridharan and Venkatappa Rao (1975)	-	306	44	262	not specified
6	Samarasinghe et al. (1982)	-	118	46	72	Ca
7	Sridharan et al. (1986a)	100	495	49.2	445.8	Na
8	Sridharan et al. (1986a)	100	233	57.8	175.2	K
9	Sridharan et al. (1986a)	100	125	40.6	84.4	Ca
10	Sridharan et al. (1986a)	100	675	49.1	625.9	Li
11	Sridharan et al. (1986b)	37	84	42	42	Ca
12	Sridharan et al. (1986b)	31	100	45.2	54.8	Ca
13	Sridharan et al. (1986b)	42	106.4	44.1	62.3	Ca
14	Sridharan et al. (1986b)	39	124.2	23.2	101	Ca
15	Wasti and Bezirci (1986)	88	526	38	488	not specified
16	Sivapullaiah et al. (1987)	-	337	55.6	281.4	not specified
17	Rao et al. (1989)	100	348	43.9	304.1	not specified
18	Acar and Olivieri (1989)	12	88	54	34	Ca
19	Acar and Olivieri (1989)	80	425	58	367	Na
20	Abdul et al. (1990)	-	470	45	425	Na
21	Di Maio and Fenelli (1994)	100	330.6	55.2	275.4	not specified
22	Abu-Hassanein et al. (1996)	85	608	43	565	Na
23	Abu-Hassanein et al. (1996)	94	516	51	465	Na
24	Di Maio (1996)	80	400	80	320	Na
25	Gleason et al. (1997)	-	603	36	567	Na
26	Gleason et al. (1997)	-	590	37	553	Na
27	Gleason et al. (1997)	-	124	26	98	Ca
28	Gleason et al. (1997)	-	123	38	85	Ca

29	Sridharan et al. (1997)	38	74.7	20	54.7	not specified
30	Petrov et al. (1997)	-	530	50	480	Na
31	Robinson and Allam (1998)	98	321	58	263	Na
32	Hettiaratchi et al. (1999)	-	407	105	302	Na
33	Hettiaratchi et al. (1999)	-	98	61	37	Ca
34	Sridharan and Nagaraj (1999)	100	320	56.6	263.4	Na
35	Kayabal and Bulus (2000)	-	320	50	270	Na
36	Karunaratne et al. (2001)	78	465	41	424	not specified
37	Tuncan et al. (2003)	88	447	60	387	Na
38	Mesri and Cepeda-Diaz (1986)/Polidori (2003)	100	205	35	170	not specified
39	Mesri and Cepeda-Diaz (1986)/Polidori (2003)	90	184.5	31.5	153	not specified
40	Mesri and Cepeda-Diaz (1986)/Polidori (2003)	80	164	28	136	not specified
41	Mesri and Cepeda-Diaz (1986)/Polidori (2003)	70	143.5	24.5	119	not specified
42	Mesri and Cepeda-Diaz (1986)/Polidori (2003)	60	123	21	102	not specified
43	Mesri and Cepeda-Diaz (1986)/Polidori (2003)	50	102.5	17.5	85	not specified
44	Young Jo et al. (2004)	88	504	39	465	Na
45	Sivapullaiah and Lakshmikantha (2004)	-	310	49	261	Na
46	Kaya and Fang (2005)	-	440	70	370	Na
47	Mishra et al. (2005)	61.4	310.5	54.1	256.4	not specified
48	Spagnoli et al. (2010)	85	455	70	385	Na
49	Shariatmadari et al. (2011)	-	199.4	41.5	157.9	not specified
50	Younus and Sreedeeep (2012)	49	224	31	193	not specified
51	Kumar Pal and Ghosh (2013)	64	159	37	122	not specified
52	Tiwari and Ajmera (2014)	-	148	49	99	not specified
53	Mir and Sridharan (2014)	63	84	25	59	Na
54	Ghazi (2015)	-	310	56	254	Na
55	Rageena and Rani (2015)	73	245	46	199	Ca
56	Ghadyani et al. (2016)	65	238	66	172	Na
58	Jang and Santamarina (2016)	-	276	44	232	not specified

57	Fan et al. (2017)	-	331.4	88.2	243.2	Ca
59	Deka and Sekharan (2017)	64	300	53	247	not specified

93 Table 1 Atterberg limits for smectitic clays with water as fluid

94

No.	Reference	< 2 μ m (%)	LL (%)	PL (%)	PI (%)
1	Andrews et al. (1967)	-	62	33	29
2	Mesri and Olson (1971)	47	45	29	16
3	Sridharan et al. (1973)	54	49	29	20
4	Sridharan and Venkatappa Rao (1975)	-	49	29	20
5	Littleton (1976)	-	83	30	53
6	Genevois (1977)	67	65	36	29
7	Genevois (1977)	67	73	44	29
8	Genevois (1977)	67	70	42	28
9	Genevois (1977)	67	69	39	30
10	Genevois (1977)	67	60	38	22
11	Genevois (1977)	67	61	39	22
12	Genevois (1977)	67	58	38	20
13	Genevois (1977)	67	64	44	20
14	Genevois (1977)	67	63	38	25
15	Genevois (1977)	67	58	37	21
16	Genevois (1977)	67	72	42	30
17	Genevois (1977)	67	122	63	59
18	Rao and Sridharan (1985)	54	49	29	20
19	Mesri and Cepeda-Diaz (1986)/Polidori (2003)	100	45	29	16
20	Mesri and Cepeda-Diaz (1986)/Polidori (2003)	90	40.5	26.1	14.4
21	Mesri and Cepeda-Diaz (1986)/Polidori (2003)	80	36	23.2	12.8
22	Mesri and Cepeda-Diaz (1986)/Polidori (2003)	70	31.5	20.3	11.2
23	Mesri and Cepeda-Diaz (1986)/Polidori (2003)	60	27	17.4	9.6
24	Mesri and Cepeda-Diaz (1986)/Polidori (2003)	50	22.5	14.5	8
25	Foreman and Daniel (1986)	98	54	31	23
26	Bowders and Daniel (1987)	-	58	34	24

27	Sivapullaiah et al. (1987)	-	51	34	17
28	Sridharan et al. (1988)	25	25	13.8	11.2
29	Sridharan et al. (1988)	27	38	15.3	22.7
30	Acar and Olivieri (1989)	90	64	34	30
31	Abdul et al. (1990)	-	61	37	24
32	Meegoda and Ratnaweera (1994)	84	48	36	12
33	Di Maio and Fenelli (1994)	100	57.5	37.8	19.7
34	Robinson and Allam (1998)	-	53	32	21
35	Sridharan and Nagaraj (1999)	11.5	58.7	45.2	13.5
36	Sridharan and Nagaraj (1999)	32	55	31.4	23.6
37	Kumar and Muir Wood (1999)	95	80	39	41
38	Lemos and Vaughan (2000)	82	69	38	31
39	Karunaratne et al. (2001)	87	74	34	40
40	Kaya and Fang (2005)	-	42	29	13
41	Sentenac et al. (2006)	-	54	31	23
42	Polidori (2007)	97	62	36	26
43	Di Matteo et al. (2011)	34	56.51	34	22.51
44	Spagnoli et al. (2012)	48	57	35	22
45	Khosravi et al. (2013)	20	45	26	19
46	Estabragh et al. (2014)	25	47	20	27
47	Tiwari and Ajmera (2014)	-	65	36	29
48	Tiwari and Ajmera (2014)	-	70	30	40
49	Ghadyani et al. (2016)	30	32	22	10
50	Jang and Santamarina (2016)	-	67	31	36
51	Fan et al. (2016)	-	29.1	19.5	9.6

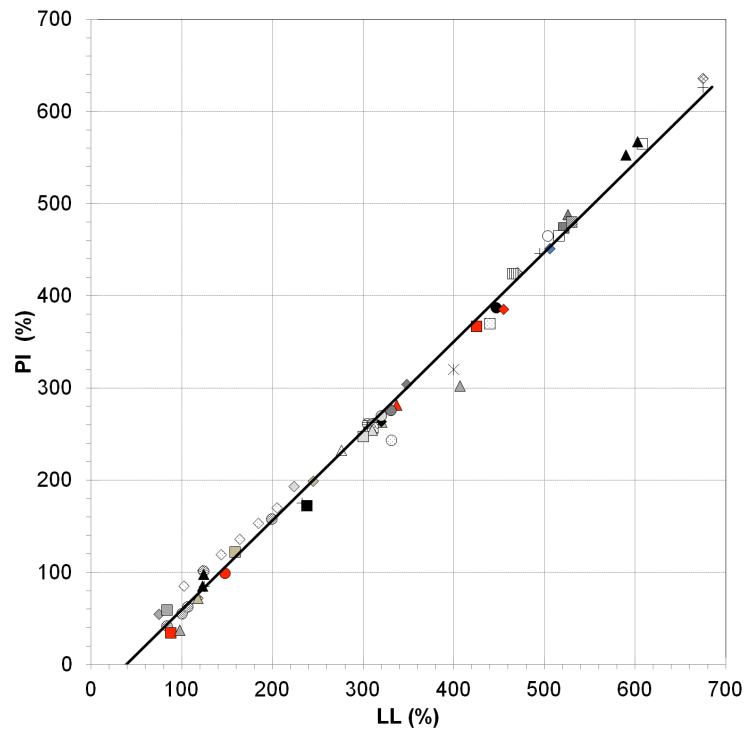
96 Table. 2 Atterberg limits for kaolinitic clays with water as fluid

97 **RESULTS AND DISCUSSION**

98 *Correlations found*

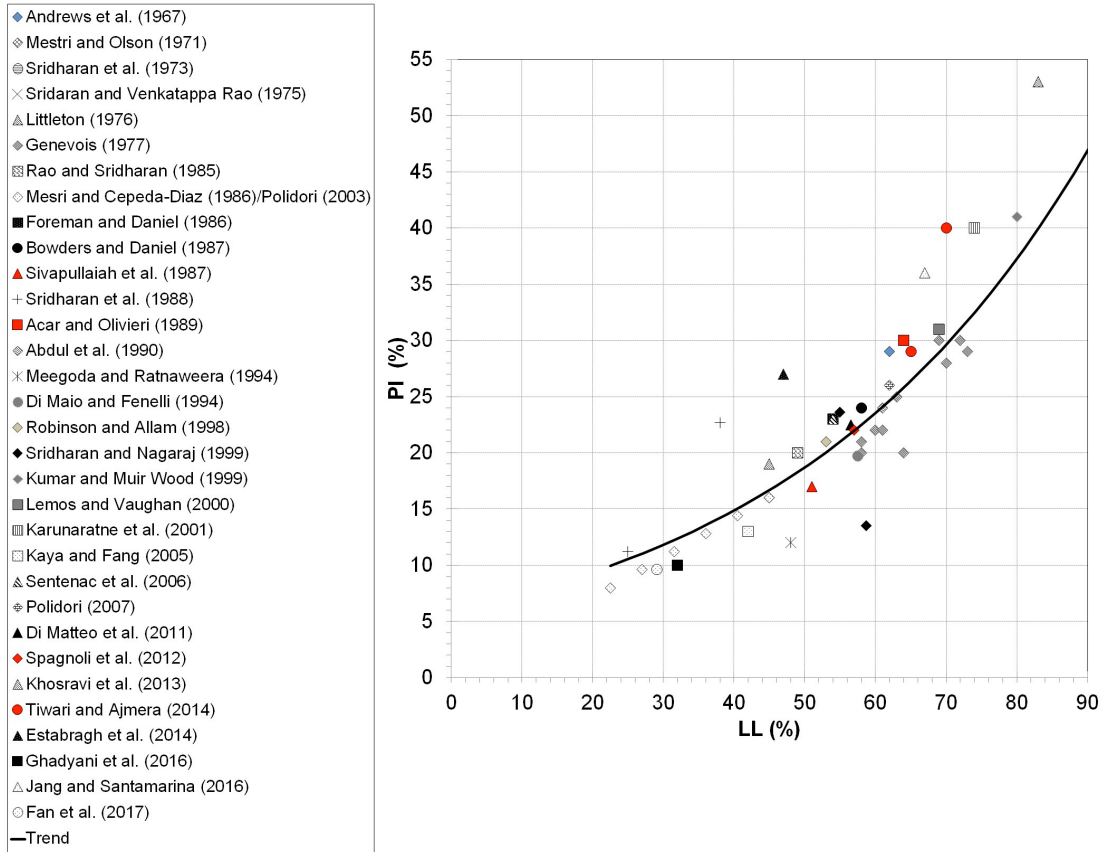
99 As LL value of clays depends on the type of clay mineral with associated cations (Mitchell
100 and Soga, 2005), smectite and kaolinite have been analyzed separately. Fig. 1 shows the
101 relation PI vs LL for smectitic clays. A linear correlation shows that $PI = 0.97 \times LL - 37.6$,
102 with a very good correlation coefficient of $R^2 = 0.99$. The correlation PI vs LL for smectite is
103 statistical significant as p-value is <0.05 (i.e. $5.18 \cdot 10^{-61}$). The correlation matches very well
104 with that found by Seed et al. (1964), where $PI = 0.98 \times LL - 27.5$, who investigated artificial
105 kaolinite-quartz mixtures in different amounts. The correlation found by Nagaraj & Jayadeva
106 (1983) was $PI = 0.74 \times LL - 8$; however, this was based on natural clays coming from different
107 depths with inhomogeneous mineralogy. Regarding the kaolinite data, an exponential
108 correlation between LL and PI was found (Fig. 2). The equation was in this case $PI =$
109 $5.94e^{0.023 \cdot LL}$ and was characterized by a correlation coefficient of $R^2 = 0.80$. With respect to
110 the data for the smectite, the results are more scattered. However, the correlation PI vs LL for
111 kaolinite is also statistical significant as p value <0.05 (i.e. $1.74 \cdot 10^{-19}$).

- Seed et al. (1964)
- ◆ Andrews et al. (1967)
- ◇ Mestri and Olson (1971)
- ⊙ Sridharan et al. (1973)
- × Sridharan and Venkatappa Rao (1975)
- ⊠ Samarasinghe et al. (1982)
- ▲ Wasti and Bezirci (1986)
- ◇ Mesri and Cepeda-Diaz (1986)/Polidori (2003)
- + Sridharan et al. (1986a)
- ⊙ Sridharan et al. (1986b)
- ▲ Sivapullaiah et al. (1987)
- ◆ Rao et al. (1989)
- Acar and Olivieri (1989)
- ◇ Abdul et al. (1990)
- Di Maio and Fenelli (1994)
- Abu-Hassanein et al. (1996)
- × Di Maio (1996)
- ▲ Gleason et al. (1997)
- ◇ Sridharan et al. (1997)
- Petrov et al. (1997)
- ▲ Robinson and Allam (1998)
- ◆ Sridharan and Nagaraj (1999)
- ▲ Hettiaratchi et al. (1999)
- Kayabal and Bulus (2000)
- Karunaratne et al. (2001)
- Tuncan et al. (2003)
- Jo et al. (2004)
- Sivapullaiah and Lakshminantha (2004)
- Mishra et al. (2005)
- Kaya and Fang (2005)
- ◆ Spagnoli et al. (2010)
- Shariatmadari et al. (2011)
- ◇ Younus and Sreedeeep (2012)
- Kumar Pal and Ghosh (2013)
- Mir and Sridharan (2014)
- Tiwari and Ajmera (2014)
- ▲ Ghazi (2015)
- ◆ Rageena and Rani (2015)
- Ghadyani et al. (2016)
- ▲ Jang and Santamarina (2016)
- Fan et al. (2017)
- Deka and Sekharan (2017)
- Trend



112

113 **Fig. 1 PI vs LL for smectitic clays.**



114

115 **Fig. 2 PI vs LL for kaolinitic clays**

116 Bearing in mind these findings, and considering the coefficient of correlations observed for
 117 smectite ($R^2 = 0.99$) and kaolinite ($R^2 = 0.80$), a comparison between experimental PI, which
 118 were derived from another published data shown in Tab. 3 and predicted PI values (obtained
 119 from the equations mentioned above) has been shown in Fig 3A & 3B. As for Tabs. 1 and 2,
 120 for the values showed in Tab. 3 only data where Atterberg limit for the Casagrande cup on
 121 pure kaolinite and smectite were selected. Smectitic soils show a linear relation with a very
 122 good R^2 value (0.99), where $PI_{\text{predicted}}=1.04 \times PI_{\text{experimental}}$.

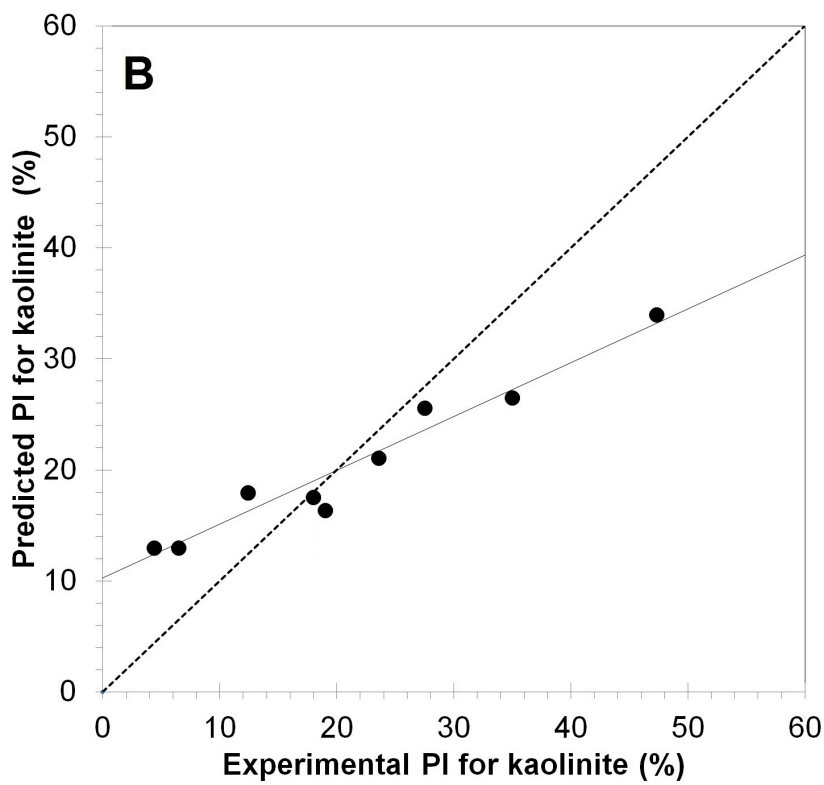
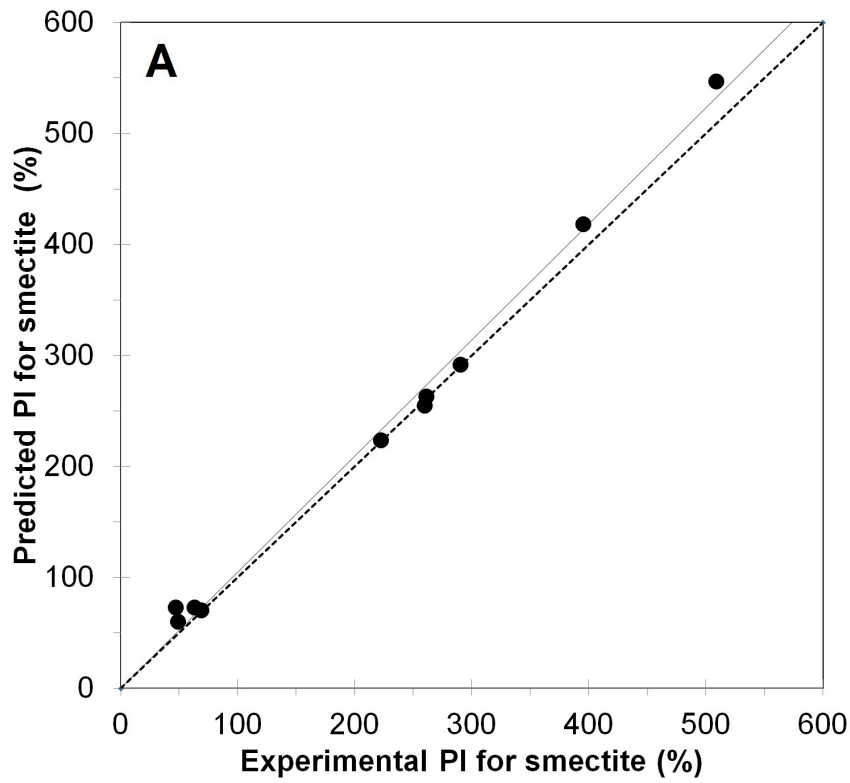
123 For kaolinitic soils the predicted PI tends to overestimate the lab. PI up to 20%. From this
 124 point forward the predicted PI values underestimate the lab PI. The relation has the form of
 125 $PI_{\text{predicted}}=0.48 \times PI_{\text{experimental}}+10.26$.

126 This is likely due to the smaller R^2 values for the PI vs LL correlation with respect to the one
 127 obtained for smectite. However, the regression coefficient, R^2 , gives a value of 0.94. Besides,
 128 both p-values for Figs 3A & 3B show also a statistical significance ($p < 0.05$) between the
 129 predicted vs lab PI values, with $3.60 \cdot 10^{-10}$ and $3.46 \cdot 10^{-06}$ for smectite and kaolinite
 130 respectively.

No.	Source	LL (%)	PI (%)	Clay type
1	Stadtbäumer (1976)	269	222	Smectite
2	Egashira and Ohtsubo (1982)	114	63	Smectite
3	Bell (1994)	114	47	Smectite
4	Stavridakis (1999)	111	68	Smectite
5	Sivapullaiah et al. (2003)	310	261	Smectite
6	Eisazadeh et al. (2012)	301	260	Smectite
7	Fatahi et al. (2013)	340	290	Smectite
8	Prakash and Sridharan (2013)	100.8	19	Smectite
9	Kolay and Ramesh (2016)	603	508	Smectite
10	Basmenj et al. (2017)	470	395	Smectite
11	Stadtbäumer (1976)	63	36	Kaolinite
12	Stavridakis (1999)	34	29	Kaolinite
13	Sridharan and Prakash (2001)	48	35	Kaolinite
14	Sridharan and Prakash (2001)	44	25	Kaolinite
15	Sridharan and Nagaraj (2005)	48	35	Kaolinite

16	Sridharan and Nagaraj (2005)	55	31	Kaolinite
17	Park et al. (2006)	47	29	Kaolinite
18	Sachan et al. (2013)	65	30	Kaolinite
19	Pulat et al. (2014)	34	27	Kaolinite
20	Kolay and Ramesh (2016)	76	28	Kaolinite

131 Table. 3 Atterberg limits for smectitic and kaolinitic clays used for the validation of the
132 predicted IP shown in Figs. 1 and 2



133

134 **Fig. 3 Predicted PI vs experimental PI for smectitic clays (A); predicted PI vs**

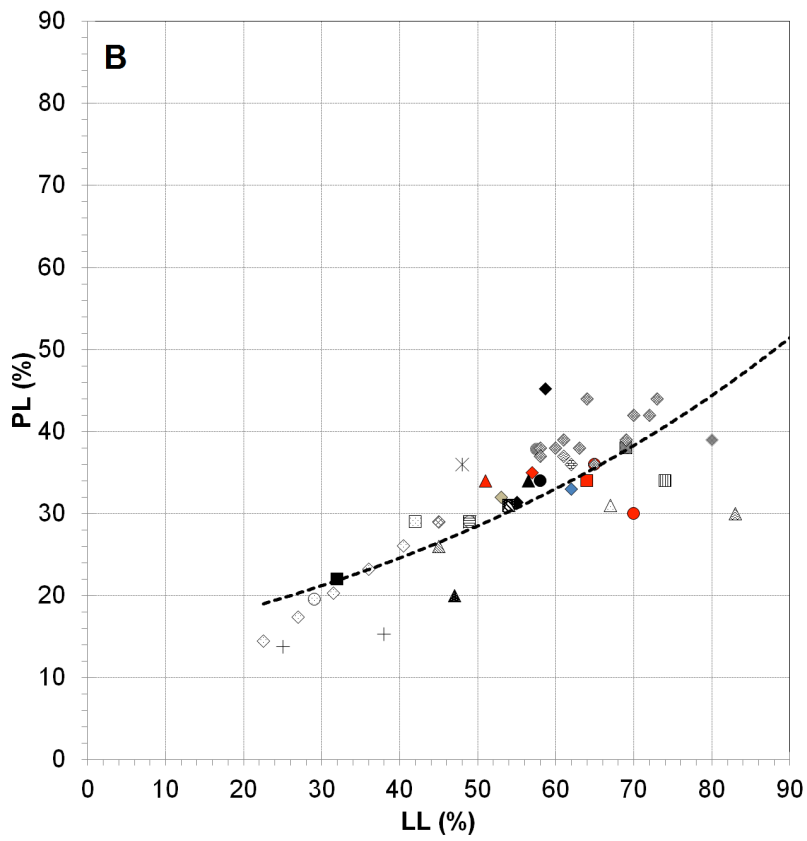
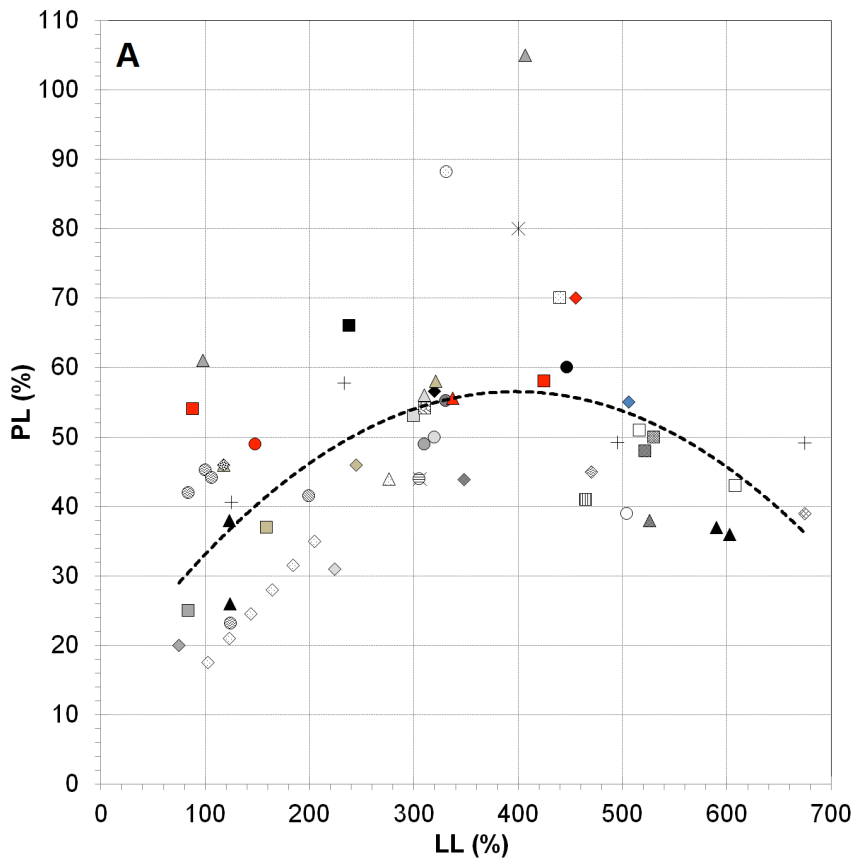
135 **experimental PI for kaolinitic clays (B). The experimental PI values refer to Tab. 3.**

136 No significant correlation between LL and the percentage of clay size fraction was found for
137 both smectite and kaolinite. This agrees with the findings of Sridharan et al. (1988), whereas
138 disagrees with the statement of Seed et al. (1964) and Polidori (2007), who presented a linear
139 variation of LL with the percentage of clay size fraction for quartz and pure clay mixtures.
140 However, it is worth mentioning that the data of the literature reviewed used in this research
141 refer to pure clays, which are normally characterized by a wider particle size distribution.

142 Fig. 4 (A and B) shows the relation PL vs LL as from Tab. 1 and 2 for both pure clays. The
143 purpose of the diagram is not to find out a relation; it is rather to show how the parameters
144 change with respect to each other. Smectitic clays (Fig. 4A) show a bell-shaped behavior,
145 where the highest PL value does not correspond to the highest LL value. The increases in PL
146 values follow increases in LL values up to a certain point, after which LL values increases
147 but PL values decreases. While LL values are directly proportional to the water content and
148 to the main cation involved, PL values show considerable variations (Bain, 1971). According
149 to Haigh et al. (2013), PL relates to the capillary suction at which the water phase ceases to
150 act as a continuum.

151 It is interesting to note that some Ca-smectites have PL values higher than the Na-smectite
152 samples. PL variations might be due to the difficulties of the thread-rolling tests and also
153 because due to the different drying (shrinkage) characteristics of the smectitic clays (Bain,
154 1971), where the shrinkage is directly proportional to the PI (Taylor and Smith, 1986).
155 Recent work shows the electrochemical forces play role in shrinkage processes (Lu and Dong,
156 2017). In that case, PL is also dependent on the electrochemical forces similar to the LL.

157 Regarding the correlation PL vs LL for kaolinitic clays (Fig. 4B), the trend is similar as
158 observed in Fig. 2, i.e. an exponential function links in an acceptable way the two parameters
159 ($R^2 = 0.70$). However, the correlations shown in Fig. 4 are not meant to be statistically
160 relevant.



162 **Fig. 4 Correlation PL vs LL for smectitic (A) and kaolinitic clays (B). The correlating**
163 **lines are dotted because they are not meant to give a statistical reference. Note that the**
164 **legend is the same as per Figs. 1 and 2.**

165 *Estimation of the statistical variability of the PI vs LL correlation*

166 In order to use the correlations obtained on the experimental measurements of Fig. 1 and 2
167 for smectite and kaolinite, an accurate probabilistic analysis is required. Since the collected
168 data show some variability regarding the estimation of PI from LL, the estimation that can be
169 made on PI leads to a probable range of variability rather than a simple deterministic value.

170 The confidence interval indicates the range that, with a certain probability (the confidence
171 level), gives the true value of the parameter (Spagnoli et al. 2017).

172 A probabilistic analysis has been assessed for the validation of correlations PI vs LL shown
173 in Fig. 1 & 2. A straight line, $y = m \cdot x + n$, better approximates the data (x_i and y_i) shown in
174 Tab. 1 and 2 and Fig. 1 and 2. The coefficients m and n were determined with Cramer's
175 method, which is useful for solving a system of linear equations using the determinant, in
176 case the system has exactly one solution. It is assumed that the error in the determination of
177 the parameter x is much smaller than one would have in the estimation of the parameter y ,
178 and that the error in the determination of each y_i is constant.

179 The uncertainty (standard deviation) σ_y on the parameters y_i is given by the following
180 equation (Bacon, 1953):

$$181 \quad \sigma_y = \sqrt{\frac{\sum_{i=1}^k (y_i - \bar{m} \cdot x_i - \bar{n})^2}{(k-2)}} \quad (1)$$

182 where k is the number of the available experimental points; \bar{m} and \bar{n} are the mean values of
183 m and n and are obtained using the equations of the linear regression on the available data.

184 Applying the theory of propagation of errors, the uncertainties of the regression coefficients
185 are obtained:

$$186 \quad \sigma_m = \sqrt{\frac{\sum_{i=1}^k (y_i - \bar{m} \cdot x_i - \bar{n})^2}{(k-2) \cdot \left(\sum_{i=1}^k x_i^2\right) - \frac{(k-2)}{k} \cdot \left(\sum_{i=1}^k x_i\right)^2}} \quad (2)$$

$$187 \quad \sigma_n = \sqrt{\frac{\left(\sum_{i=1}^k x_i^2\right) \cdot \sum_{i=1}^k (y_i - \bar{m} \cdot x_i - \bar{n})^2}{k \cdot (k-2) \cdot \left(\sum_{i=1}^k x_i^2\right) - (k-2) \cdot \left(\sum_{i=1}^k x_i\right)^2}} \quad (3)$$

188 The best estimation of the parameters m and n in relation to a certain level of confidence
 189 (expressed in percentage) is obtained through the confidence intervals:

$$190 \quad m = \bar{m} \pm t_{\eta\%,k-2} \cdot \sigma_m \quad (4a)$$

$$191 \quad n = \bar{n} \pm t_{\eta\%,k-2} \cdot \sigma_n \quad (4b)$$

192 where $t_{\eta\%,k-2}$ is the parameter t of Student, relative to a confidence level of $\eta\%$ and a degree
 193 of freedom of $k-2$. Student's t -distribution is a distribution of continuous probability
 194 governing the relationship between two random variables, the first with normal distribution,
 195 while, the second follows a squared distribution.

196 Based on the discussion above, it is possible to predict the value y_0 and its uncertainty, σ_{y_0} , for
 197 a value of x_0 . In fact:

$$198 \quad \bar{y}_0 = \bar{m} \cdot x_0 + \bar{n} \quad (5)$$

199 Applying, then, the theory of propagation of errors to the equation 5, the standard deviation
 200 of y_0 is obtained by:

$$201 \quad \sigma_{y_0} = \sqrt{\frac{\sum_{i=1}^k (y_i - \bar{m} \cdot x_i - \bar{n})^2}{(k-2)} \cdot \left[\frac{\left[\left(\sum_{i=1}^k x_i^2 \right) - \frac{1}{k} \cdot \left(\sum_{i=1}^k x_i \right)^2 \right] + \left[k \cdot x_0 - \left(\sum_{i=1}^k x_i \right) \right]^2}{k \cdot \left(\sum_{i=1}^k x_i^2 \right) - \left(\sum_{i=1}^k x_i \right)^2} \right]} \quad (6)$$

202 The confidence interval for y_0 will be given, then, by the following expression, for a
 203 confidence level of $\eta\%$:

$$204 \quad y_0 = \bar{y}_0 \pm t_{\eta\%,k-2} \cdot \sigma_{y_0} \quad (7)$$

205 That means:

$$206 \quad y_0 = (\bar{m} \cdot x_0 + \bar{n}) \pm t_{\eta\%,k-2} \cdot \sqrt{\frac{\sum_{i=1}^k (y_i - \bar{m} \cdot x_i - \bar{n})^2}{(k-2)} \cdot \frac{\left[\left(\sum_{i=1}^k x_i^2 \right) - \frac{1}{k} \cdot \left(\sum_{i=1}^k x_i \right)^2 \right] + \left[k \cdot x_0 - \left(\sum_{i=1}^k x_i \right) \right]^2}{k \cdot \left(\sum_{i=1}^k x_i^2 \right) - \left(\sum_{i=1}^k x_i \right)^2}}$$

208 In this way, it will be possible to identify a range of variability of y , relative to a certain
 209 desired confidence level, corresponding to the parameter x_0 indicated.

210 Considering the data of this research for smectite (Fig. 1), the relation LL (x axis) vs PI (y
 211 axis) is shown. The linear regression analysis allowed to identify the line that best
 212 approximates the experimental points available (PI = 0.97 x LL - 37.6). The standard
 213 deviation on the slope "m" is $\sigma_m = 0.012$; the standard deviation on constant term "n" is $\sigma_n =$

214 4.157. Moreover, we have: $k = 59$, $\sum_{i=1}^k (y_i - \bar{m} \cdot x_i - \bar{n})^2 = 13417$, $\left(\sum_{i=1}^k x_i \right) = 18302$,

215 $\left(\sum_{i=1}^k x_i^2 \right) = 7381520$. If, for example, the confidence interval for PI with a confidence level of

216 95%, corresponding to a LL equal to 350% is requested, then PI = 301.38 ÷ 302.48. The same
 217 PI confidence interval for a greater confidence level of 99%, would be PI = 301.20 ÷ 302.66.

218 In the case of kaolinite, the relation PI -LL is exponential, as it can be seen from Fig. 2.

219 Considering the same approach explained before, it is possible to analyze the relationship
 220 ln(PI)-LL for the available experimental data. In this case, it is possible to apply the linear
 221 regression and obtain, then, the line that best approximates the experimental data available.

222 This line takes the form: $\ln(\text{PI}) = 0.026 \cdot \text{LL} + 1.61$.

223 The standard deviation on the slope "m" is $\sigma_m = 0.0012$; the standard deviation on the
 224 constant term "n" is $\sigma_n = 0.0998$. Besides, we observe: $k = 51$, $\sum_{i=1}^k (y_i - \bar{m} \cdot x_i - \bar{n})^2 = 1.52$,
 225 $\left(\sum_{i=1}^k x_i \right) = 2725.31$, $\left(\sum_{i=1}^k x_i^2 \right) = 158641$. If, for example, the confidence interval for PI with a
 226 confidence level of 95%, corresponding to a LL equal to 50% is requested, then
 227 $\ln(\text{PI}) = 2.91 \div 2.93$ and therefore $\text{PI} = 18.48 \div 18.75$. The same PI confidence interval for a
 228 greater confidence level of 99%, would be $\text{PI} = 18.44 \div 18.80$.
 229 Both probabilistic assessments clearly confirm the robustness of the regression functions
 230 obtained for this research if the data shown in Fig. 1 and 2 are considered.

231 CONCLUSIONS

232 In this paper, results obtained from 44 technical papers for smectitic and kaolinitic clays
 233 respectively, published from different countries and authored by several authors over a period
 234 of 50 years have been analyzed with respect to Atterberg limits and some useful conclusions
 235 have been drawn. Good correlations have been obtained between liquid limit (LL) and
 236 plasticity index (PI) for both smectite and kaolinite. The range of LL varied up to 680% for
 237 smectitic and 85% for kaolinitic clays. It can be shown that the correlation of LL with PI for
 238 smectitic clays varies marginally with different ranges of LL. The predicted PI values haven
 239 been verified against another independent set of data, giving very good R^2 values, namely
 240 0.99 and 0.94 for smectite and kaolinite respectively, although the predicted PI values
 241 compared with the lab PI values for kaolinite are not as precise as for smectite. The validity
 242 of the relation PI vs LL for both clay types has been also verified by a probabilistic analysis.
 243 Besides, for smectitic clays it has been seen that PL increases along with LL up to a certain
 244 value and then it tends to decrease while LL still increases. Whereas, for kaolinitic clays, PL
 245 constantly increases with increasing LL values.

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488 **FIGURE CAPTIONS**

489 Fig. 1 PI vs LL for smectitic clays.

490 Fig. 2 PI vs LL for kaolinitic clays

491 Fig. 3 Predicted PI vs experimental PI for smectitic clays (A); predicted PI vs experimental PI
492 for kaolinitic clays (B). The experimental PI values refer to Tab. 3.

493 Fig. 4 Correlation PL vs LL for smectitic (A) and kaolinitic clays (B). The correlating lines
494 are dotted because they are not meant to give a statistical reference. Note that the legend is
495 the same as per Figs. 1 and 2.